# Geochemistry of the Earth's Surface

Edited by Halldór Ármannsson Orkustofnun, Reykjavík, Iceland

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## Invited lecture: Atmospheric CO<sub>2</sub>, terrestrial ecology, and mammalian evolution

Thure E.Cerling – Department of Geology and Geophysics, University of Utah, Salt Lake City, Utah, USA

James R. Ehleringer – Department of Biology, University of Utah, Salt Lake City, Utah, USA

John Harris – George G. Page Museum, Los Angeles, Calif., USA

Bruce MacFadden – Florida Museum of Natural History, Gainesville, Fla., USA

ABSTRACT: The distribution of  $C_3$  versus  $C_4$  plants in the modern world is related to temperature. However, because of the sensitivity of  $C_3$  to the  $CO_2/O_2$  ratio of the atmosphere it is expected that global changes in the distribution of  $C_3$  versus  $C_4$  plants should accompany atmospheric  $CO_2$  changes in geological time. The period from 6 to 8 million years ago was a time of  $C_4$  expansion in many parts of the globe.

#### I INTRODUCTION

There is strong ecological, geographical, and palaeoecological evidence that abundance and distribution of plants using C<sub>4</sub> photosynthesis are related to climate, specifically to atmospheric carbon dioxide concentration and temperature. C<sub>4</sub> plants are mainly grasses, especially the tropical and subtropical grasses (Sage et al. 1999).

### 2 RESULTS AND DISCUSSION

The C<sub>1</sub> photosynthetic pathway evolved under very high atmospheric CO2 concentrations whereas the C4 photosynthetic pathway is a much more recent development, representing adaptation to relatively low atmospheric CO2 concentrations. A model of quantum yield of plants using the C3 versus C4 photosynthetic pathways shows that the crossover temperatures for C<sub>3</sub> plants compared to C<sub>4</sub> monocots is about 22°C for the modern condition (Cerling et al. 1997). This results from photorespiration in C<sub>3</sub> plants at low CO<sub>2</sub>/O<sub>2</sub> ratios. The model further shows that the crossover is strongly dependent on both temperature and atmospheric CO<sub>2</sub> (Figure 1). C4 dicots are more rare than C4 monocots and have significantly lower crossover temperatures; they are favored only showing the crossoverunder extremely low CO2 conditions, such as are found under full Glacial conditions (Ehleringer et al. 1997). The "C<sub>4</sub>world", where C4 plants make up a significant fraction of tropical to temperate ecosystems, began at the end of the Miocene due to low atmospheric CO<sub>2</sub> concentrations; since then C<sub>3</sub> plants have been starved for CO2. C4 grasses are the dominant grass from 0 to 45 degrees latitude, and thus modern savannas are important ecosystems that characterize the "C4-world" of the Plio-Pleistocene. Mammalian evolution during the Plio-Pleistocene was directly related to the low CO2 concentrations of the atmosphere, because competition between C4 and C3 plants (and even between different C<sub>3</sub> plants) resulted from CO<sub>2</sub> starvation of C<sub>3</sub> plants due to photorespiration. Some mammals, such as equids, were already grazing on C<sub>3</sub> grasses and immediately used this new resource when it became more readily available at the end of the Miocene. Figure 2 shows that equids had already apparently adapted to a grazing diet in North America by about 15 million years ago as indicated by their development of hypsodont (high crowned) teeth, but changed to a C<sub>4</sub> diet abruptly between 6 and 7 million years ago. The change to high crowned teeth in equids was during a period of diversification within the equids.

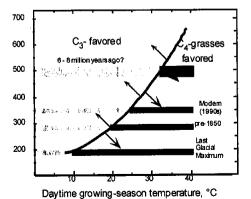


Figure 1. Model of C<sub>3</sub> versus C<sub>4</sub> photosynthesis based on relative quantum yields.

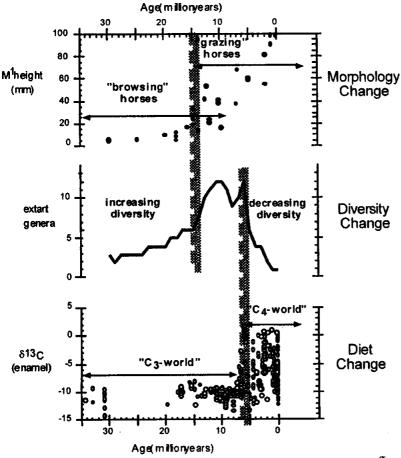


Figure 2. Index of hypsodonty for North American equids, the number of extant genera, and  $\delta^{17}$ C of equid tooth enamel from North America in the Neogene.

whereas the time of C<sub>4</sub> expansion was accompanied by a rapid decline in equid diversity (Figure 2).

The change to a C<sub>4</sub> diets in mammals was abrupt: Figure 3 shows that 10 million years ago Asian, African, and the North American equids had a C<sub>3</sub> diet, but changed to a C<sub>4</sub> diet between 6 and 8 million years ago. Entire mammalian communities on all continents (except Antarctica) underwent a great change from the "C<sub>3</sub>-world" at the end of the Miocene to the Plio-Pleistocene "C<sub>4</sub>-world", even in regions where C<sub>4</sub> plants have never been abundant (e.g., Europe).

Many vegetation changes in the Pliocene and Pleistocene have been explained as being the result of changes in seasonality or changes in aridity (e.g., Janis 1993). However, water loss from plants occurs through the stomata which is where CO<sub>2</sub> is taken up by plants through diffusion. Thus, reduction in photosynthesis and ultimately growth rate, could be

due to lower atmospheric  $CO_2$  concentrations instead of aridity. For this reason, changes in  $CO_2$  concentration also should be considered in climate reconstruction because the direct impact of  $CO_2$  on plant productivity is independent of water and temperature constraints.

#### 3 CONCLUSIONS

Global changes in terrestrial ecology are expected to result from changes in the  $CO_2/O_2$  ratio. At  $CO_2$  concentrations less than about 500 ppmV  $C_3$  plants increasingly become less efficient due to photorespiration. Therefore changes in the fraction of  $C_3$  versus  $C_4$  plants may be related to changes in the atmospheric  $CO_2$  concentration. The late Miocene to early Pliocene was a period of global vegetation change as well as a period of faunal

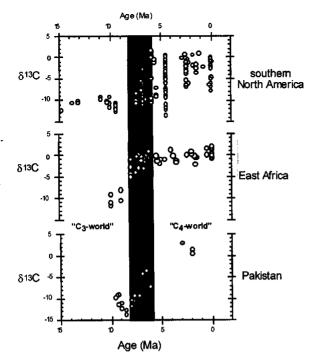


Figure 3. Global diet change in equids shown for North America, Africa, and Asia based on the  $\delta^{13}C$  values of equid tooth enamel

turnover in the mammals. The last 6 to 8 million years can be characterized as a " $C_4$ -world" which differed greatly in terrestrial ecology than the previous " $C_3$ -world".

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Nekrasov, I.Ya. 90 5410 723 5 Geochemistry, mineralogy and genesis of gold deposits

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